

THE LONG-TERM EVOLUTION OF THE GEOBIOSPHERE

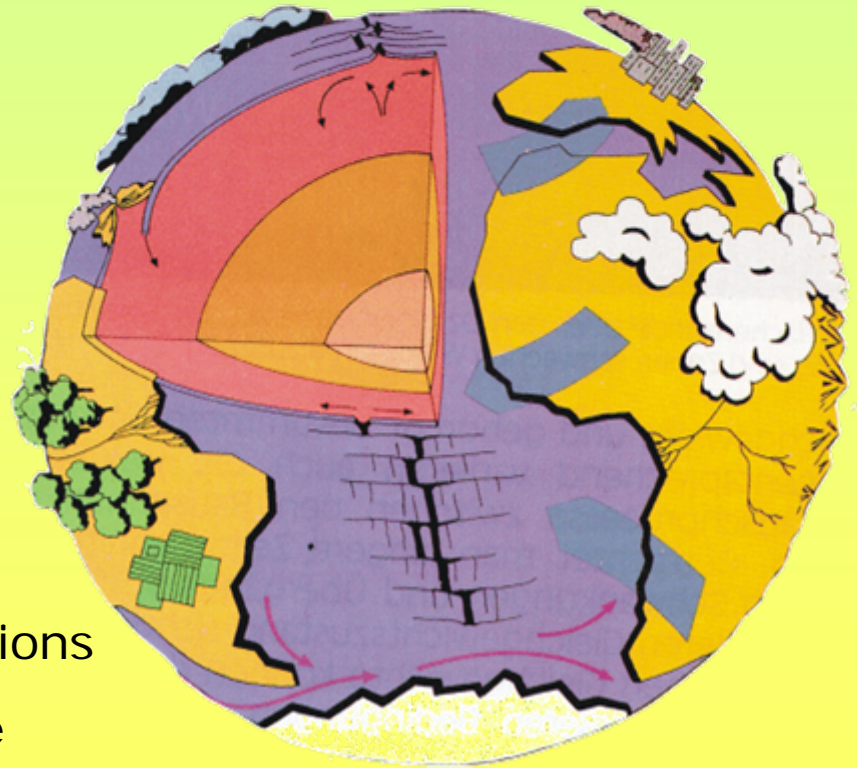


Siegfried Franck

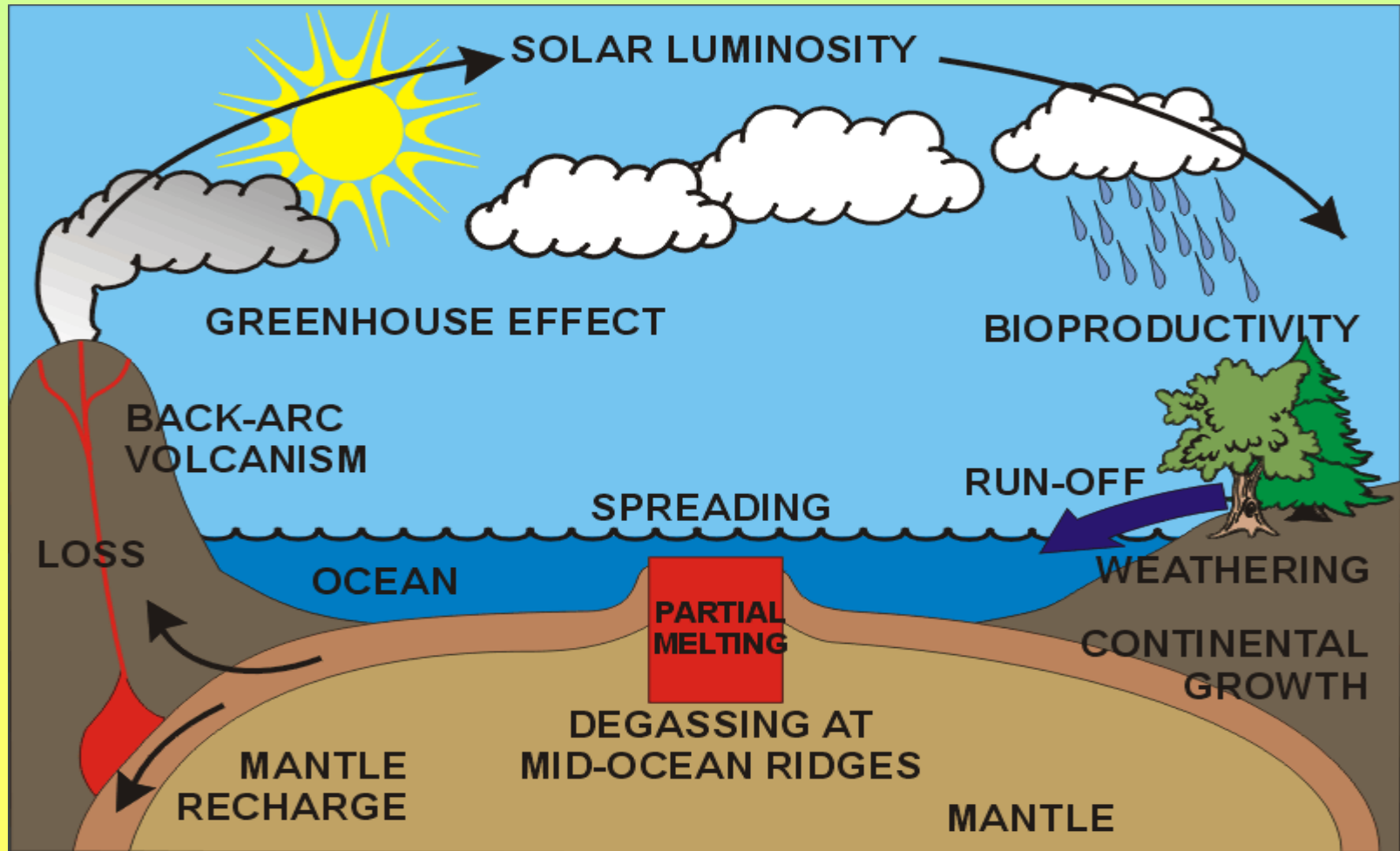


*Potsdam Institute for
Climate Impact Research*

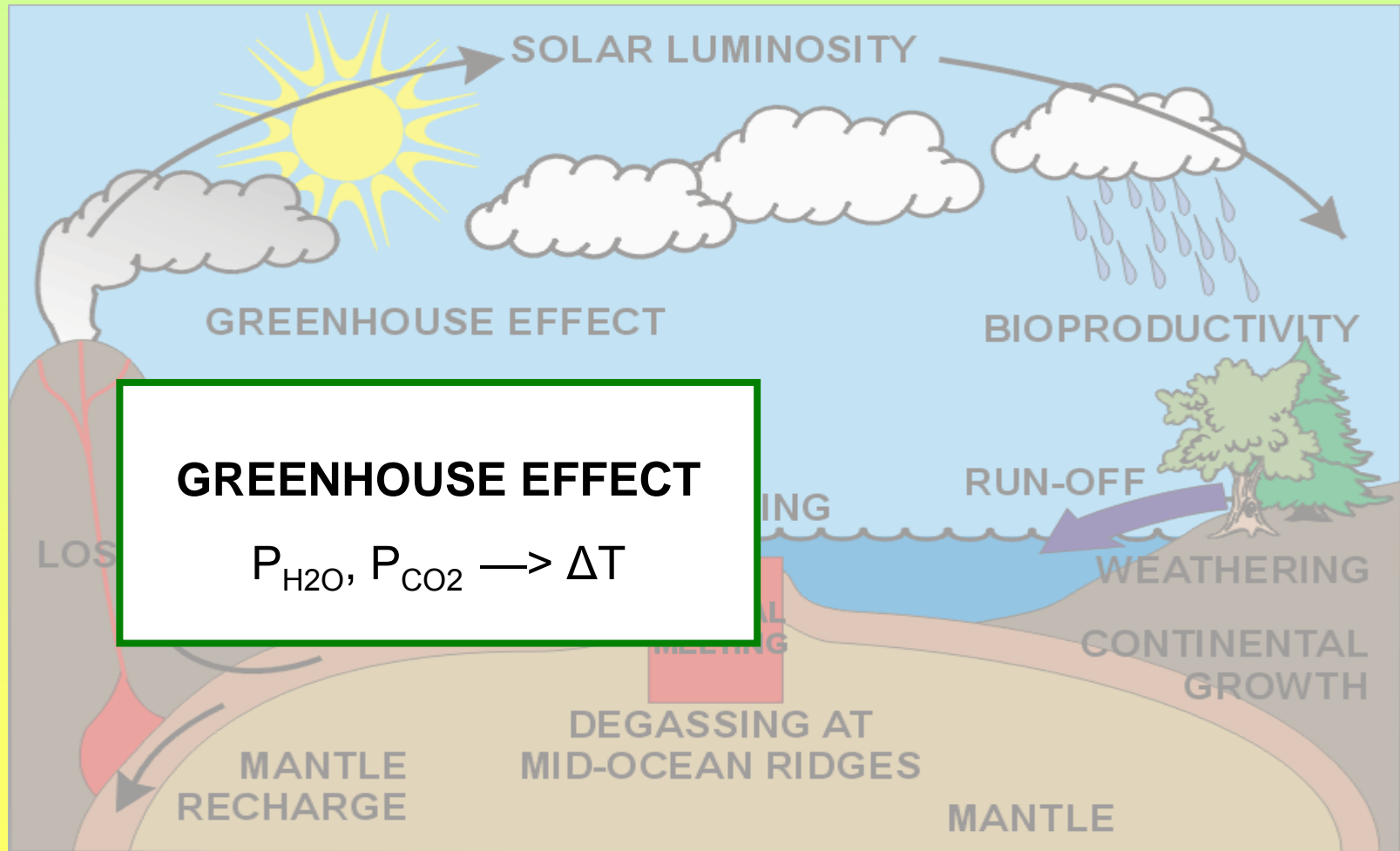
1. The global carbon cycle
2. Biosphere-geosphere interactions
3. The life span of the biosphere



THE GLOBAL CARBON CYCLE

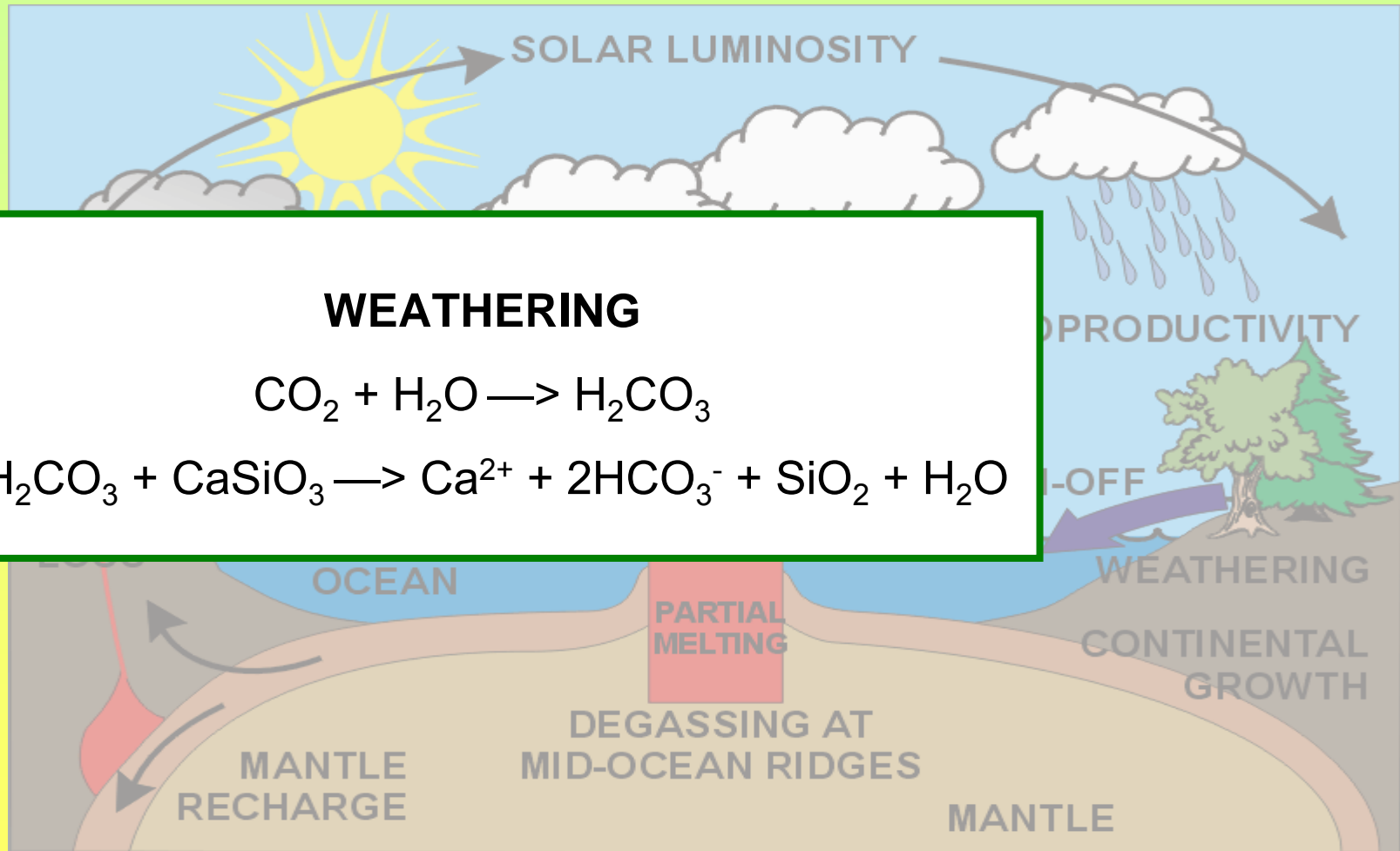
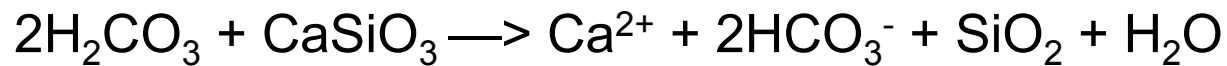
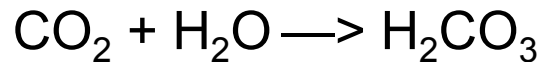


THE GLOBAL CARBON CYCLE

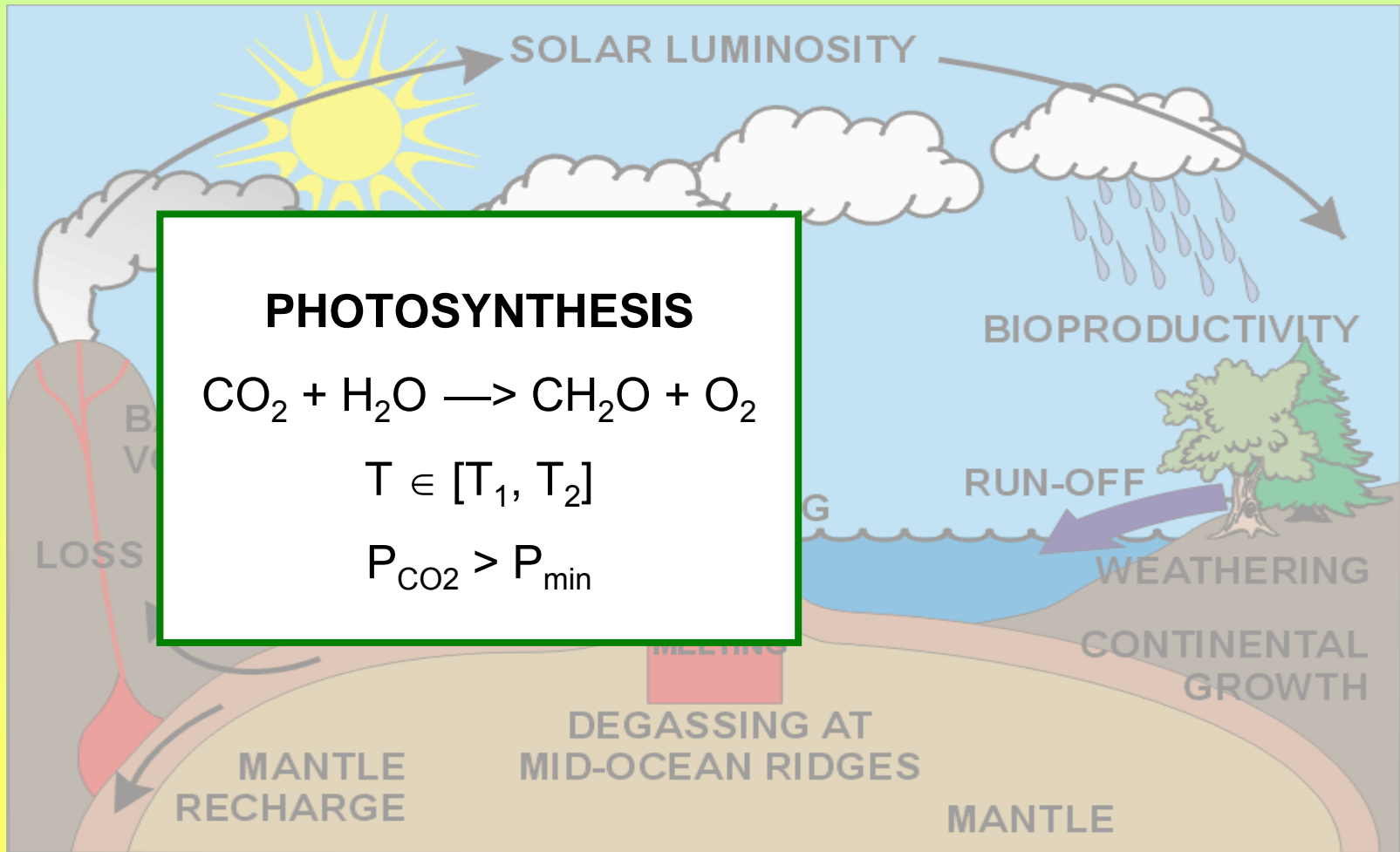


THE GLOBAL CARBON CYCLE

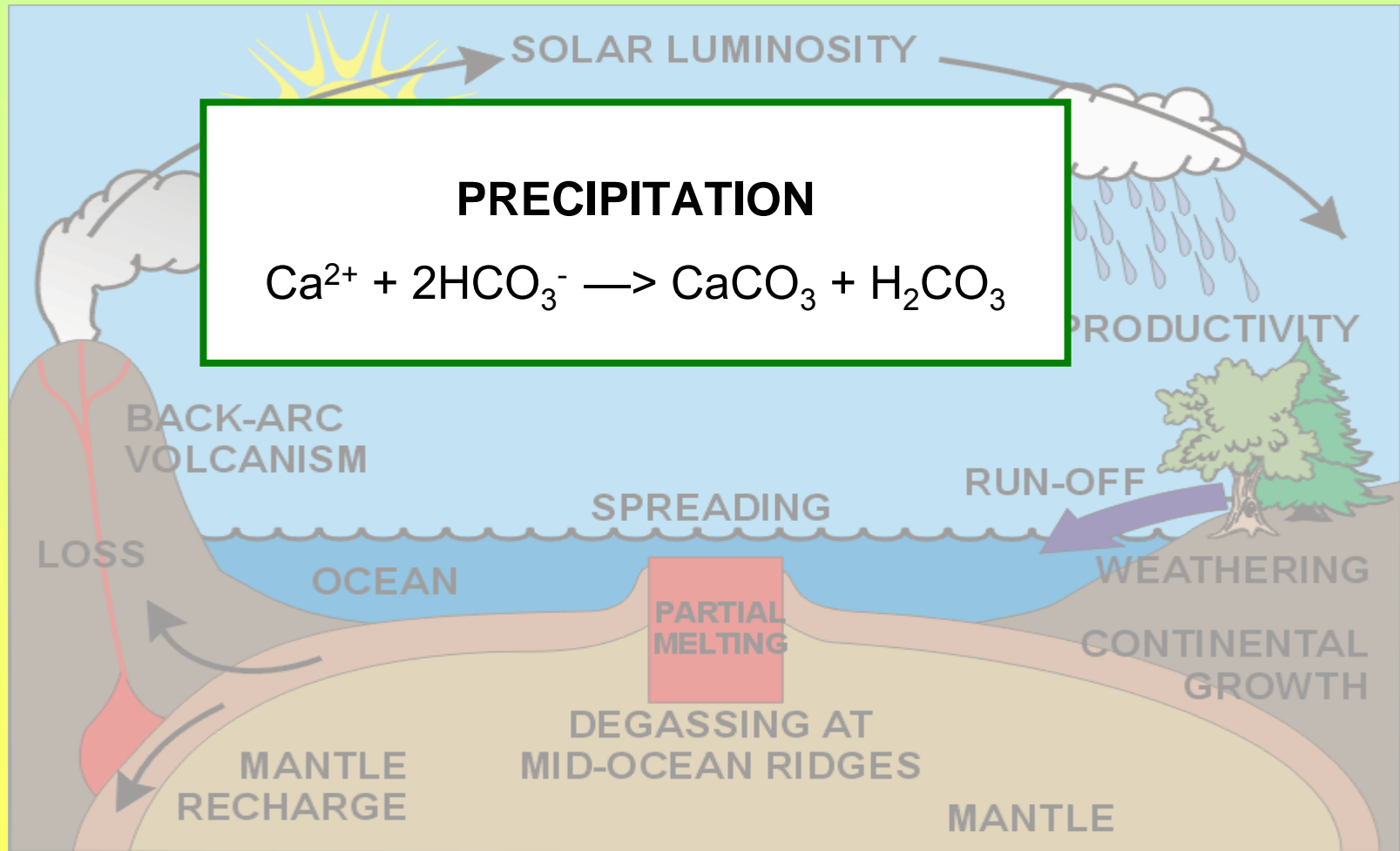
WEATHERING



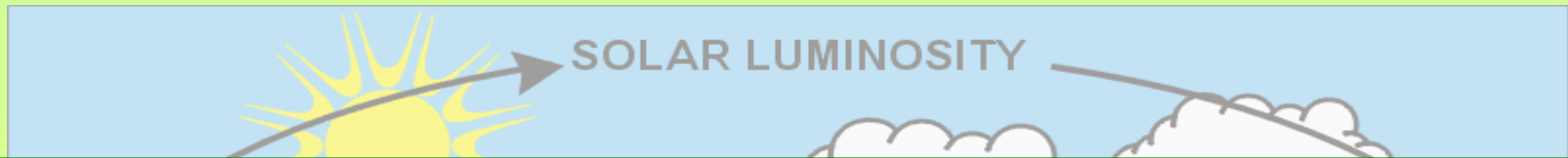
THE GLOBAL CARBON CYCLE



THE GLOBAL CARBON CYCLE

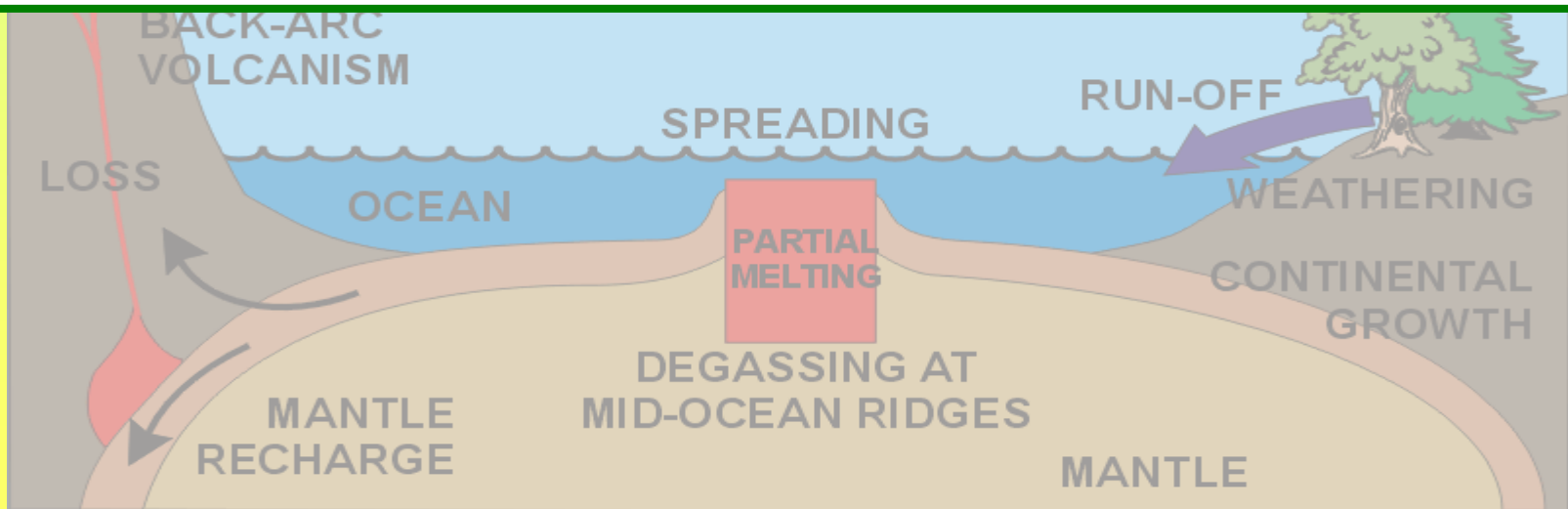


THE GLOBAL CARBON CYCLE

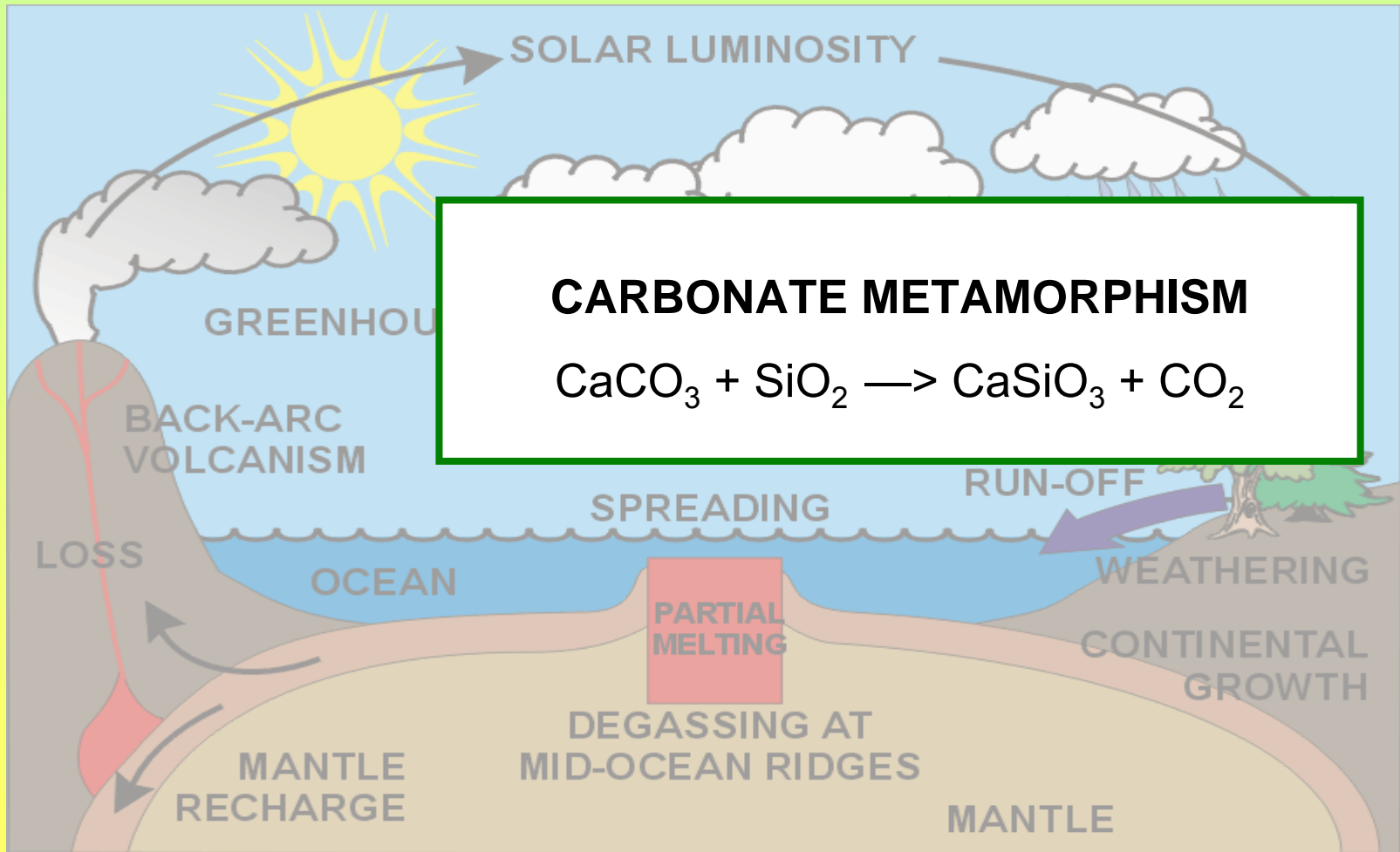


HYDROTHERMAL REACTIONS

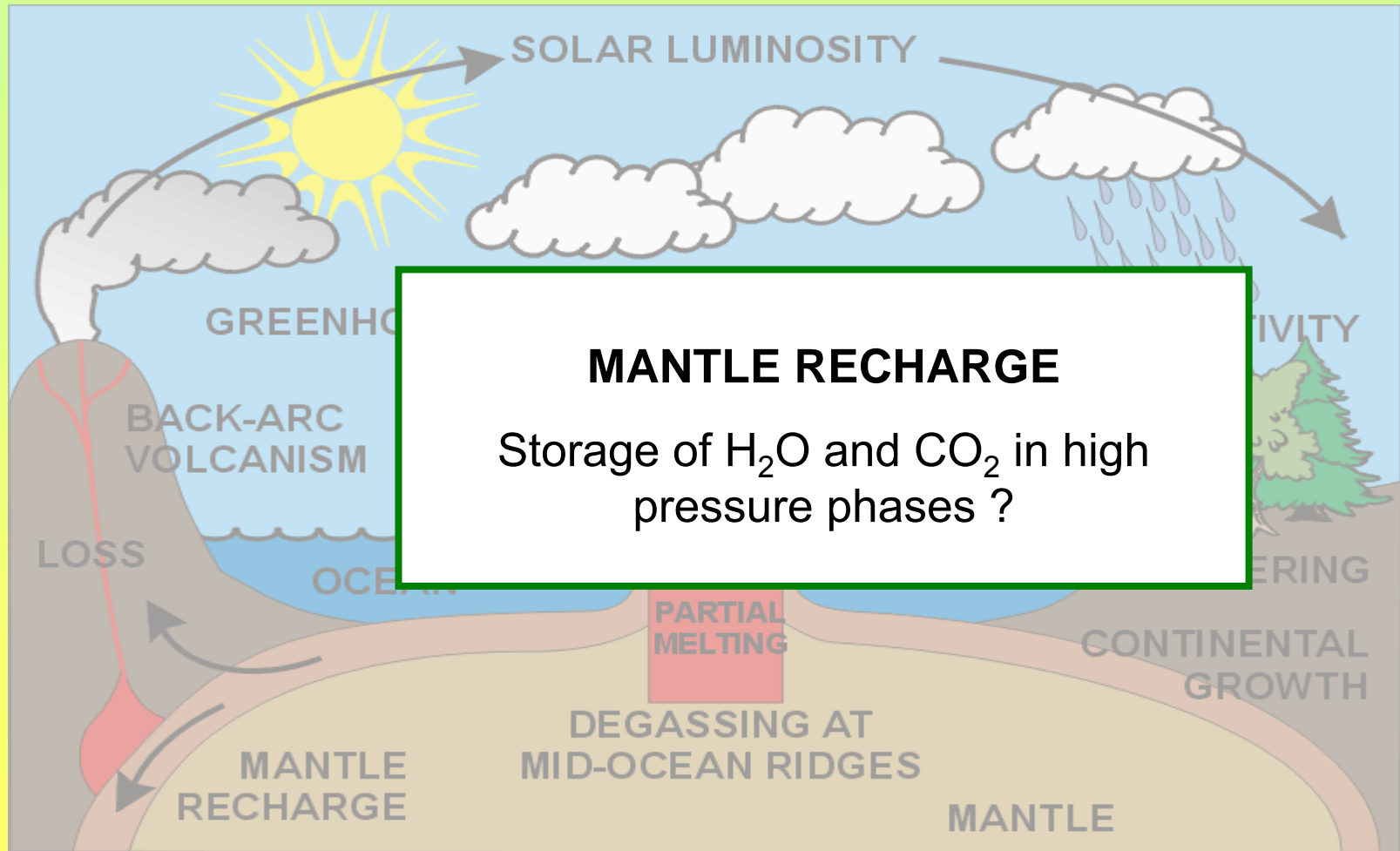
CO_2 (from seawater) + cations (from fresh basalts) \rightarrow carbonates



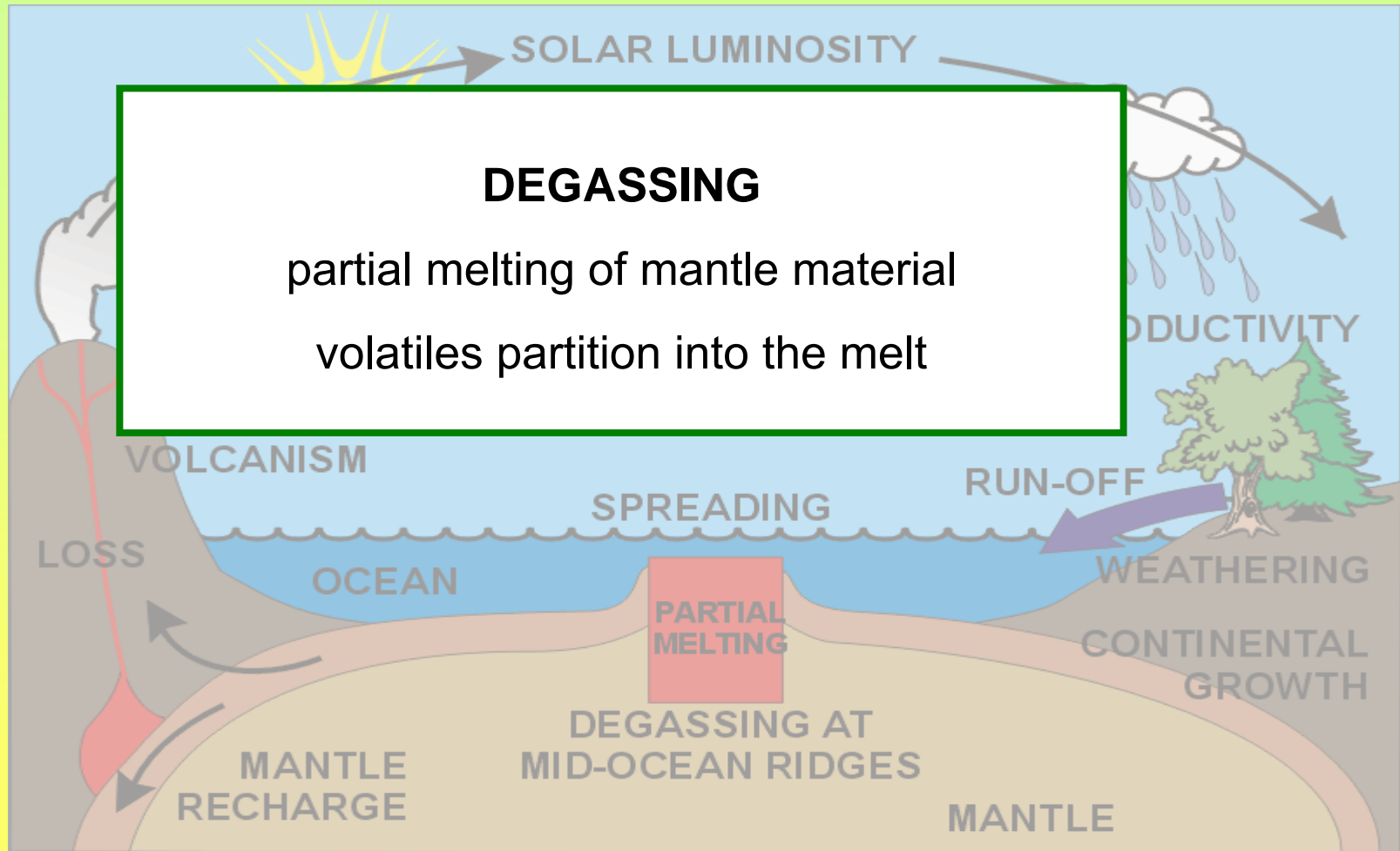
THE GLOBAL CARBON CYCLE



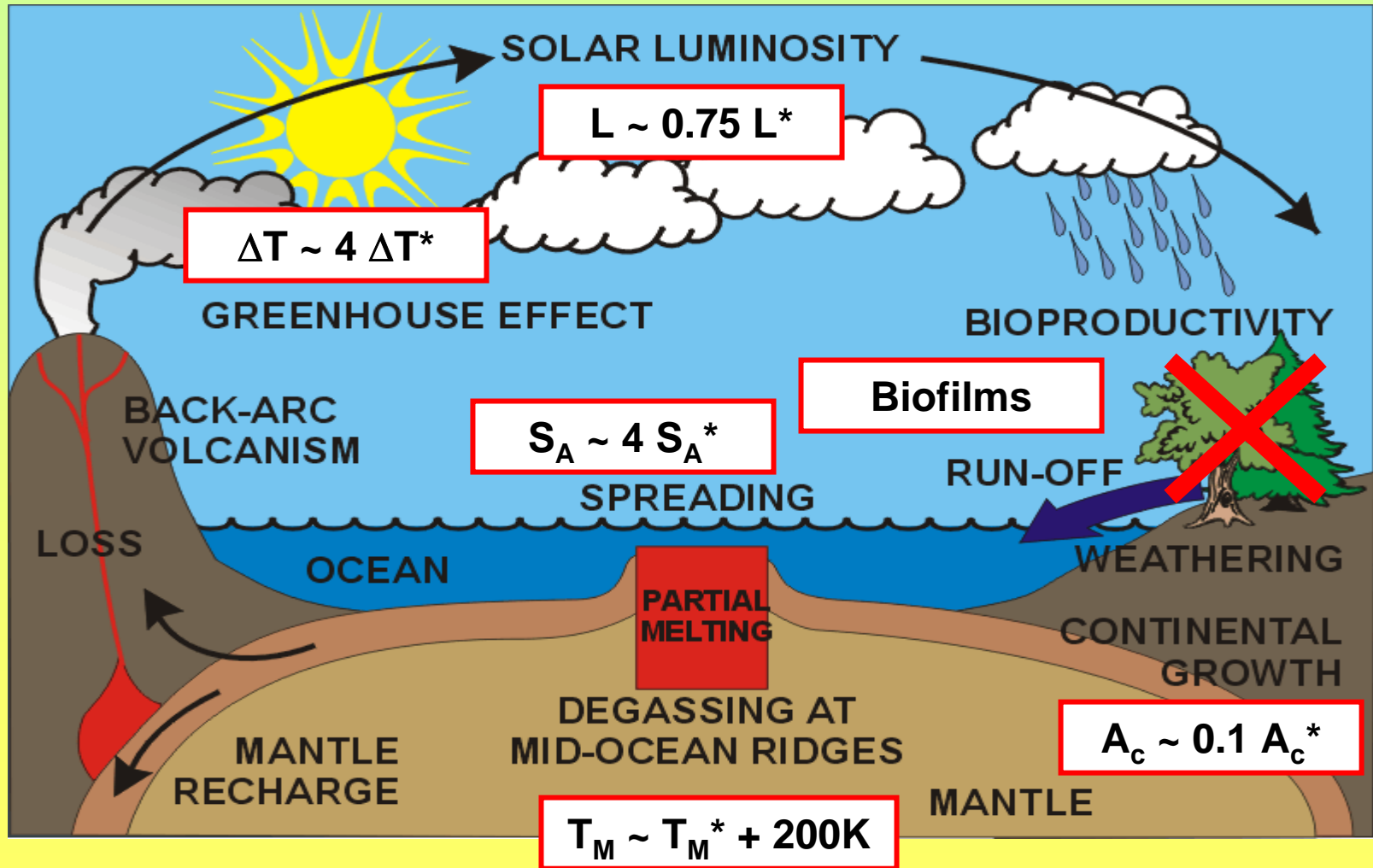
THE GLOBAL CARBON CYCLE



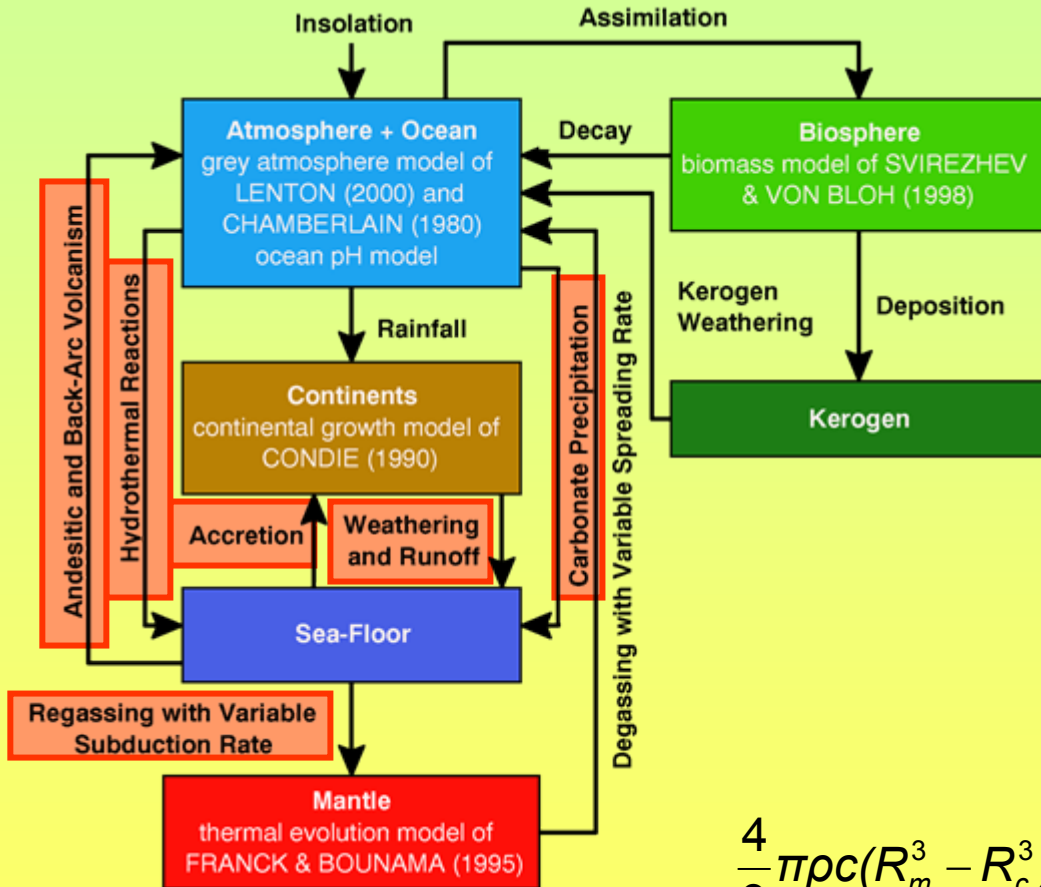
THE GLOBAL CARBON CYCLE



THE ARCHAEOAN GLOBAL CARBON CYCLE



MODEL DESCRIPTION



$$\dot{C}_m = \tau_f^{-1}(1-A)RC_f - S_A f_c d_m C_m / V_m$$

$$\begin{aligned} \dot{C}_{o+a} = & \tau_f^{-1}(1-A)(1-R)C_f \\ & + S_A f_c d_m C_m / V_m + F_{weath} \\ & + (1-\gamma)\tau_{bio}^{-1}C_{bio} + \tau_{ker}^{-1}C_{ker} \\ & - \Pi_{bio} - F_{prec} - F_{hyd} \end{aligned}$$

$$\dot{C}_c = \tau_f^{-1}AC_f - F_{weath}$$

$$\dot{C}_f = F_{prec} + F_{hyd} - \tau_f^{-1}C_f$$

$$\dot{C}_{bio} = \Pi_{bio} - \tau_{bio}^{-1}C_{bio}$$

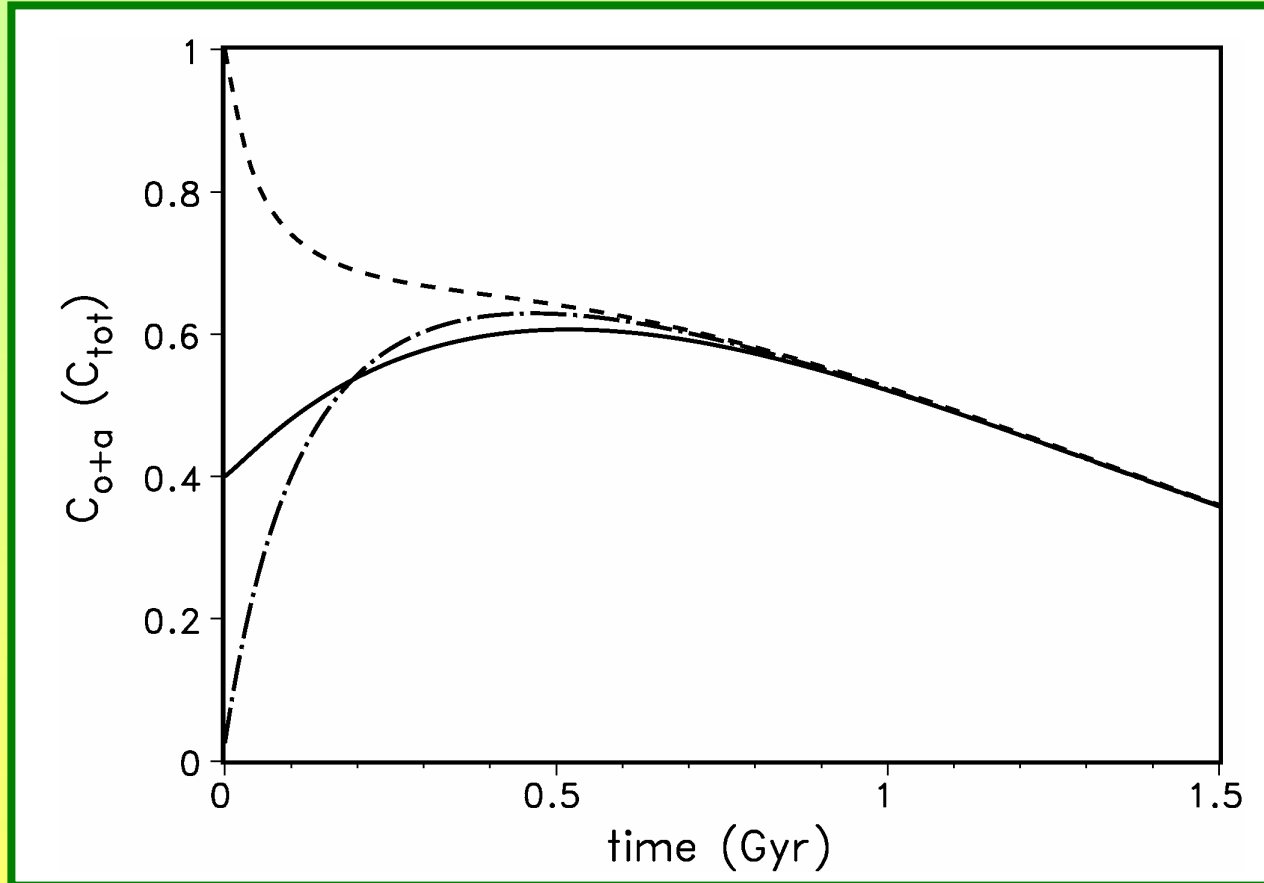
$$\dot{C}_{ker} = \gamma\tau_{bio}^{-1}C_{bio} - \tau_{ker}^{-1}C_{ker}$$

$$\frac{4}{3}\pi\rho c(R_m^3 - R_c^3)\frac{dT_m}{dt} = -4\pi R_m^2 q_m + \frac{4}{3}\pi Q(R_m^3 - R_c^3)$$

Franck et al. (2002), Tellus 54B, 325.



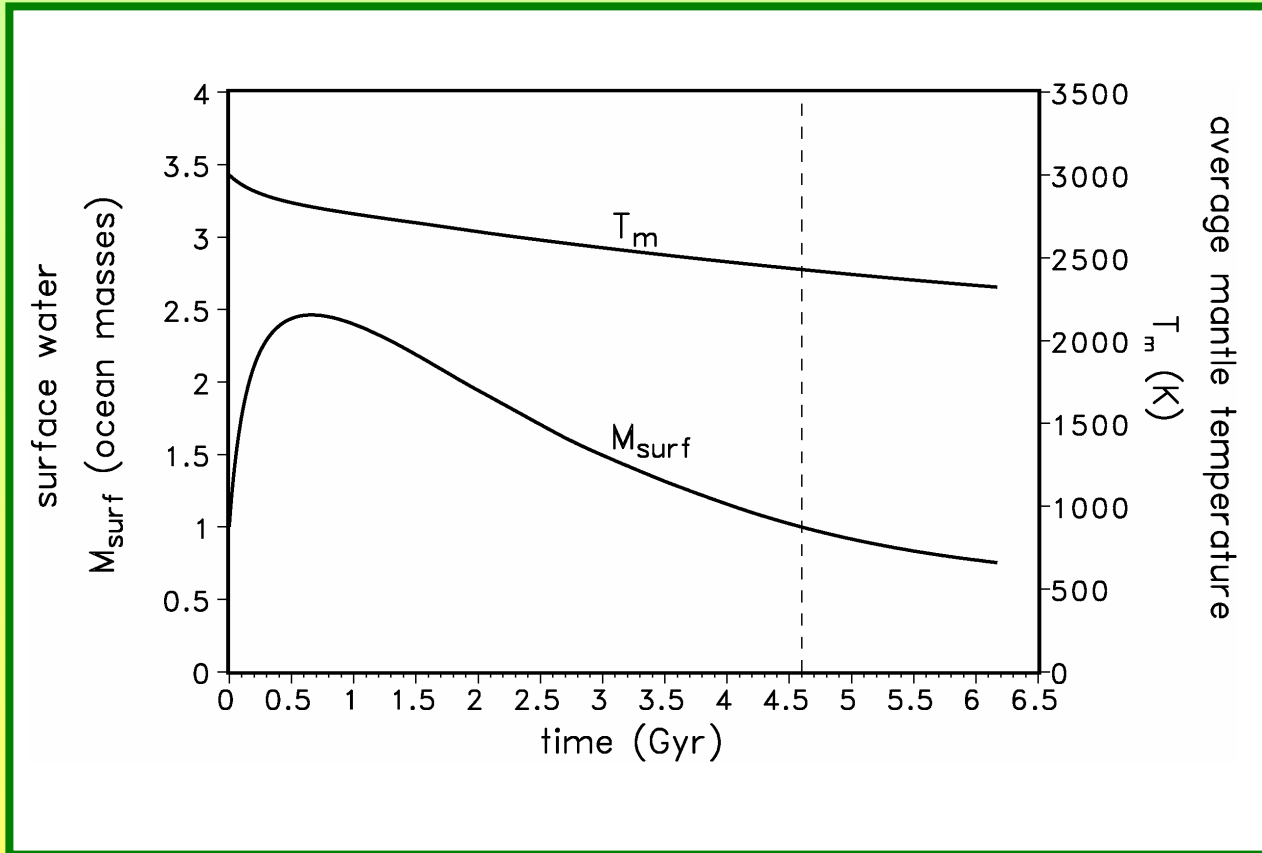
ADJUSTMENT



Evolution of the combined ocean and atmosphere reservoir C_{o+a} scaled to the total amount of carbon C_{tot} for various initial distributions of carbon between the pools. Note that after 1 Gyr the system has “forgotten” the initial conditions.



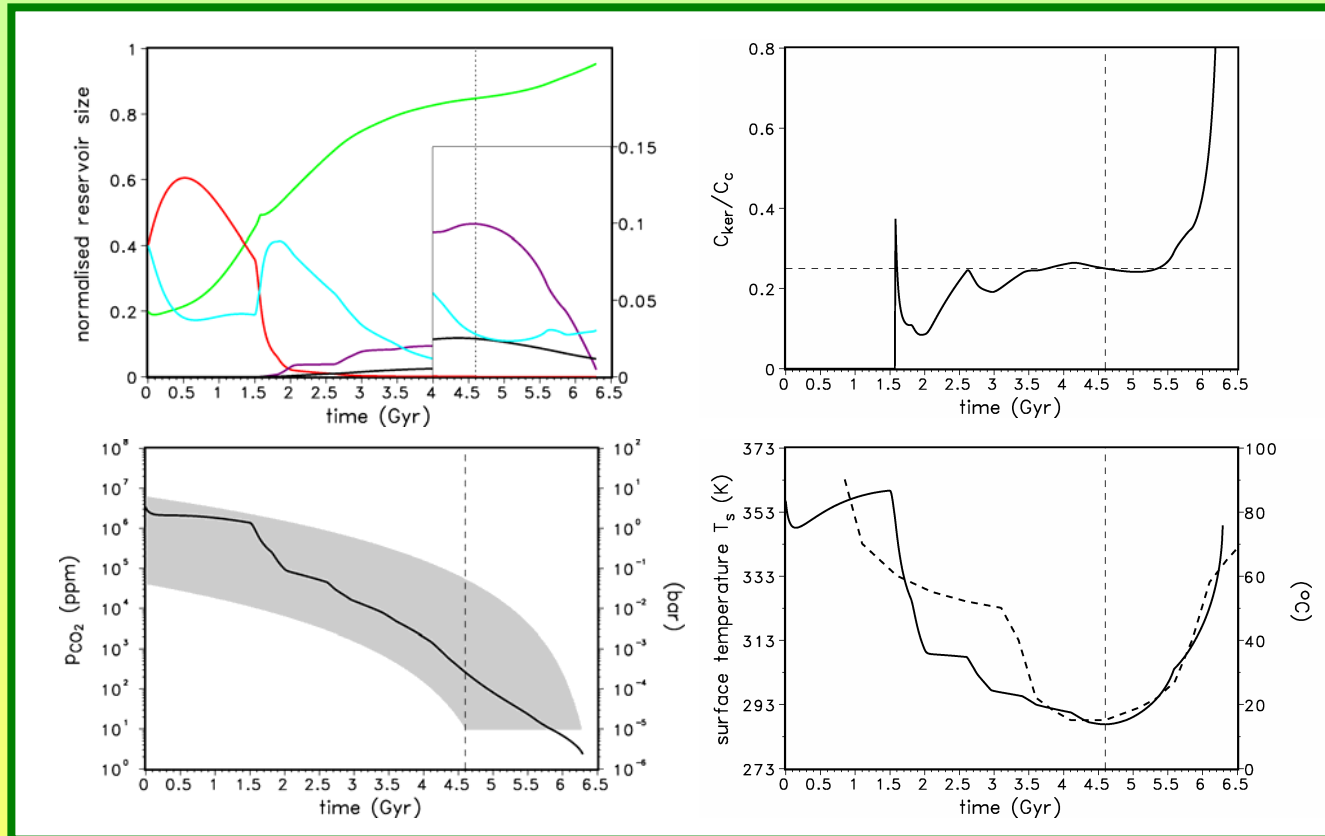
T_m and M_{surf}



Evolution of the average mantle temperature T_m and the surface water reservoir M_{surf} derived from the thermal evolution model.



RESULTS FOR CONSTANT HYDROTHERMAL FLUX

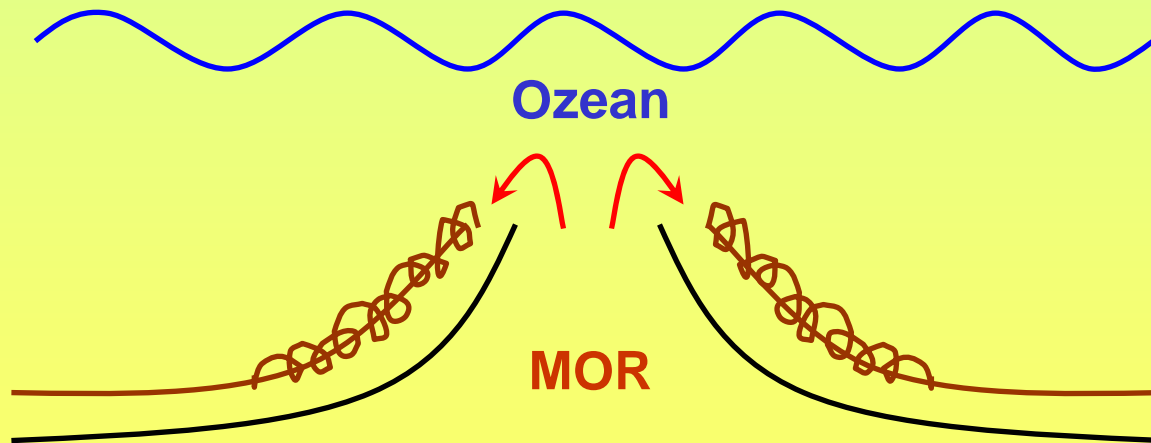


Model results for constant hydrothermal flux: (a) evolution of the reservoirs mantle (green), atmosphere + ocean (red), ocean floor (blue), kerogen (black), and continents (magenta), (b) evolution of the ratio C_{ker}/C_c where the horizontal dashed line indicates the observed ratio of 0.25, (c) evolution of atmospheric partial pressure of carbon dioxide pCO_2 where the grey shaded area represents values for non-vanishing biological productivity, and (d) evolution of the surface temperature T_s where the dashed line indicates the result of Schwartzman (1999).



HYDROTHERMAL CARBONISATION

- = Reactions of „warm“ (ca. 20°C) water with fresh basalts of the oceanic crust where carbonates are formed
- = CO₂ flux from the reservoir ocean+atmosphere to the reservoir oceanic crust



CO₂ from water + cations from basalts → carbonates

HYDROTHERMAL FLUXES

Parameterisation of hydrothermal fluxes F_{hyd}

1. $F_{\text{hyd}} = \text{constant}$
2. $F_{\text{hyd}} \propto SR$
3. $F_{\text{hyd}} \propto SR \propto p\text{CO}_2$ (Sleep and Zahnle, 2001)

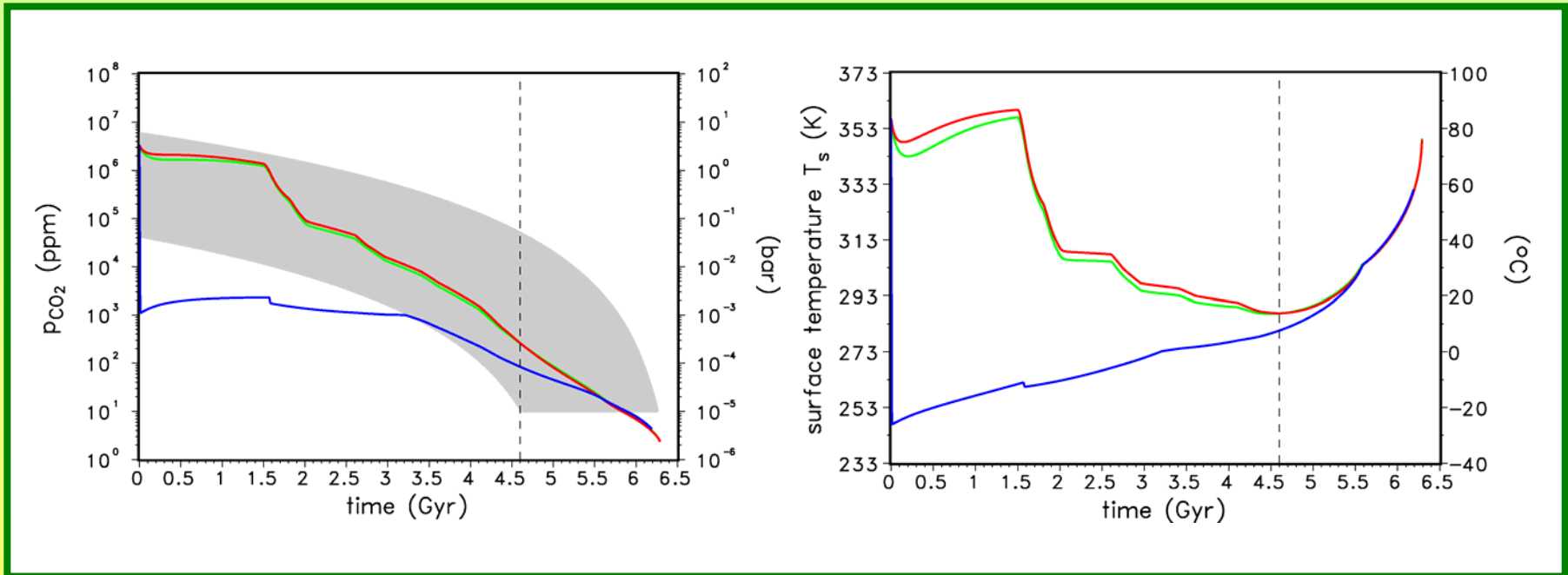
where

SR : spreading rate

$p\text{CO}_2$: partial pressure of CO_2 in the ocean



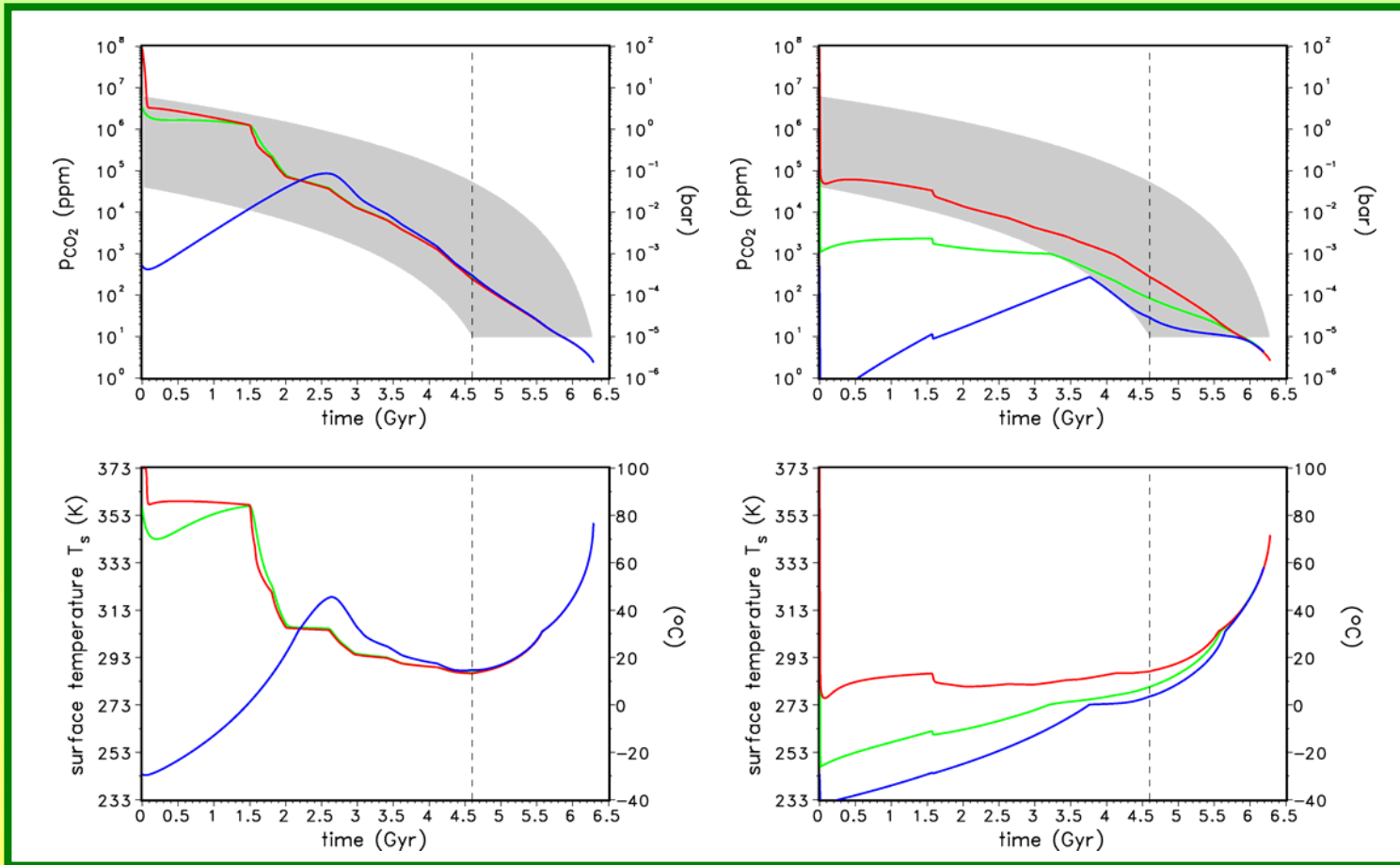
3 PARAMETERISATIONS FOR F_{hyd}



Evolution of atmospheric CO₂ partial pressure p_{CO_2} where the grey shaded area represents values for non-vanishing biological productivity (a) and surface temperature T_s (b) for three different parameterisations of the hydrothermal flux: constant (red), slow hydrothermal reaction kinetics (green), and fast hydrothermal reaction kinetics (blue).



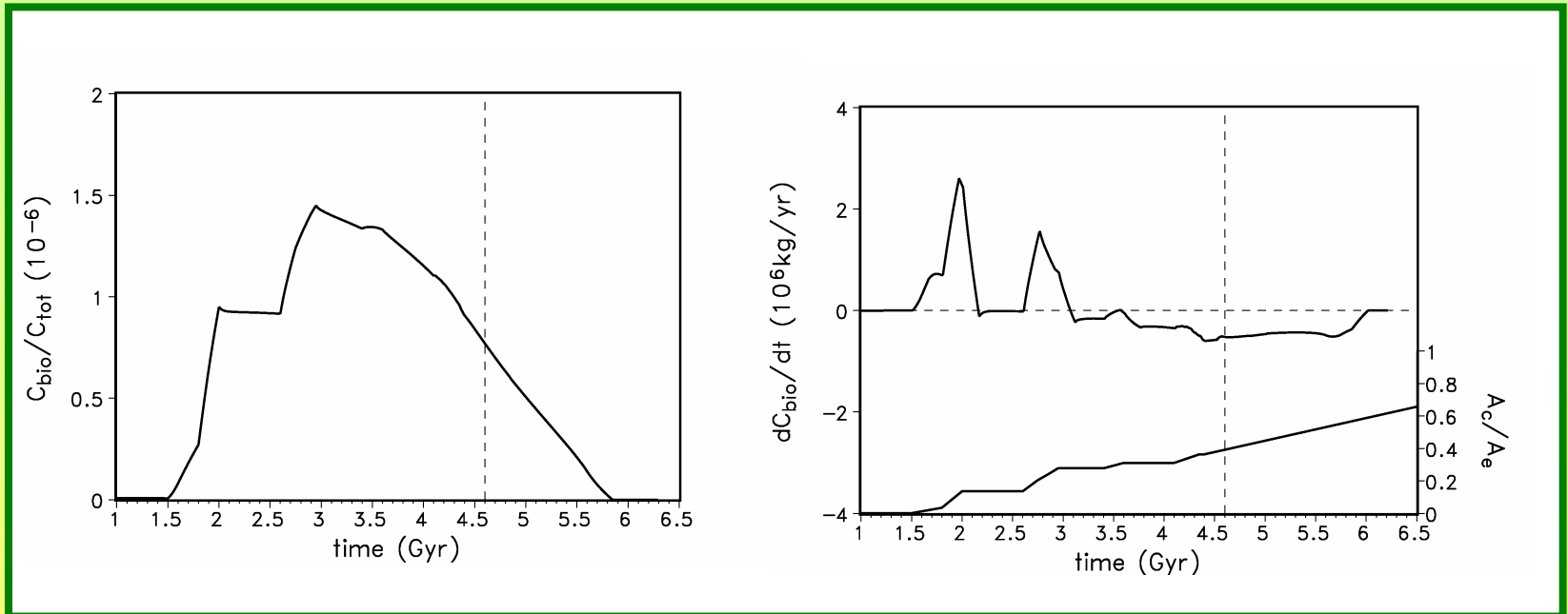
pH MODELS



Evolution of atmospheric CO₂ partial pressure p_{CO_2} where the grey shaded area represents values for non-vanishing biological productivity (a,b) and surface temperature T_s (c,d) under the condition of slow hydrothermal reaction kinetics (a,c) and fast hydrothermal reaction kinetics (b,d) for three different ocean pH models: acid ocean model (red), constant pH (green), and soda ocean model (blue)..

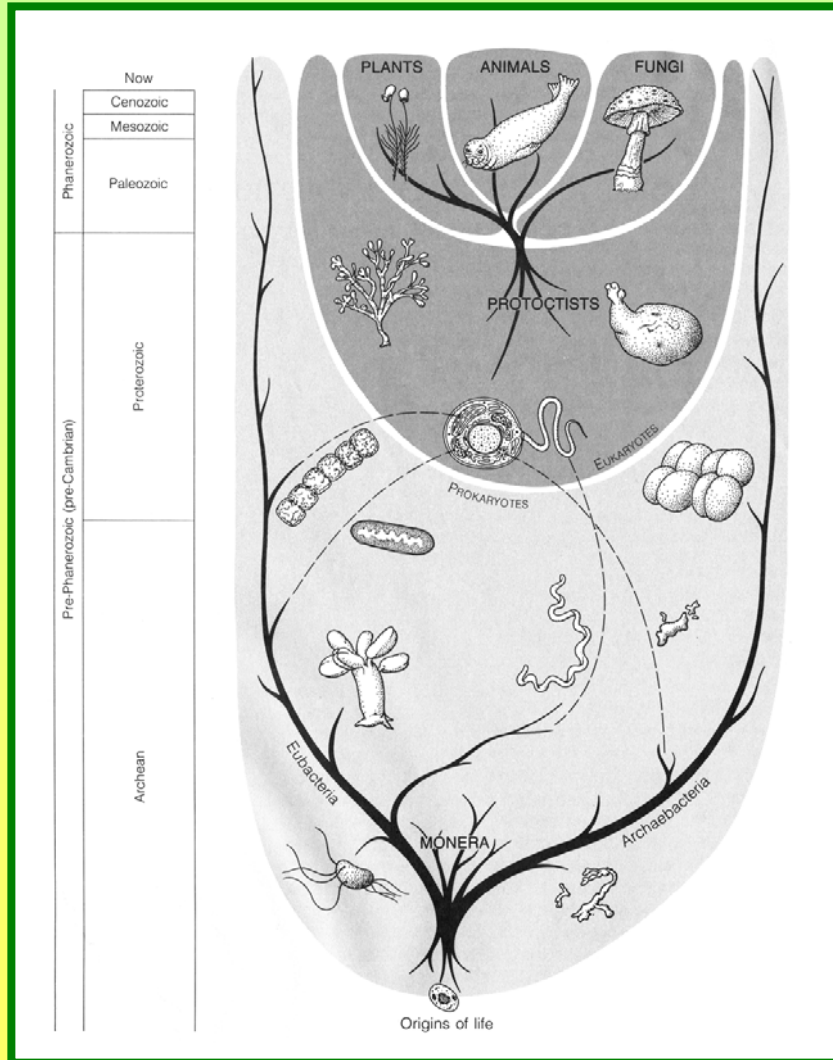


THE BIOSPHERE POOL C_{bio}



Evolution of the biosphere pool C_{bio} (a) and its time derivative dC_{bio}/dt , (b) for slow hydrothermal reaction kinetics to emphasise changes in the biosphere pool. In order to demonstrate the correlation between the continental growth rate and changes in the biosphere pool the Condie model (Condie 1990) is displayed additionally in (b).

EVOLUTION OF BIOSPHERE



Symbiosis in cell evolution

Lynn Margulis (1993)

GEOSPHERE-BIOSPHERE FEEDBACKS

Biological productivity:

$$\Pi_{\text{bio},i} = \Pi_{\text{max},i} \cdot \Pi_{T,i}(T) \cdot \Pi_{\text{pCO}_2,i}(\text{pCO}_2), i = 1,2,3$$

temperature tolerance windows		
$i=1$	procaryotes	[2°C, 100°C]
$i=2$	eucaryotes	[5°C, 45°C]
$i=3$	complex multicellular life	[0°C, 30°C]



GEOSPHERE-BIOSPHERE FEEDBACKS

Biological enhancement of weathering:

$$F_{\text{weath}} \propto \beta \cdot (a_{H^+})^{0.5} \cdot \exp \frac{T - T^*}{13.7\text{K}}$$

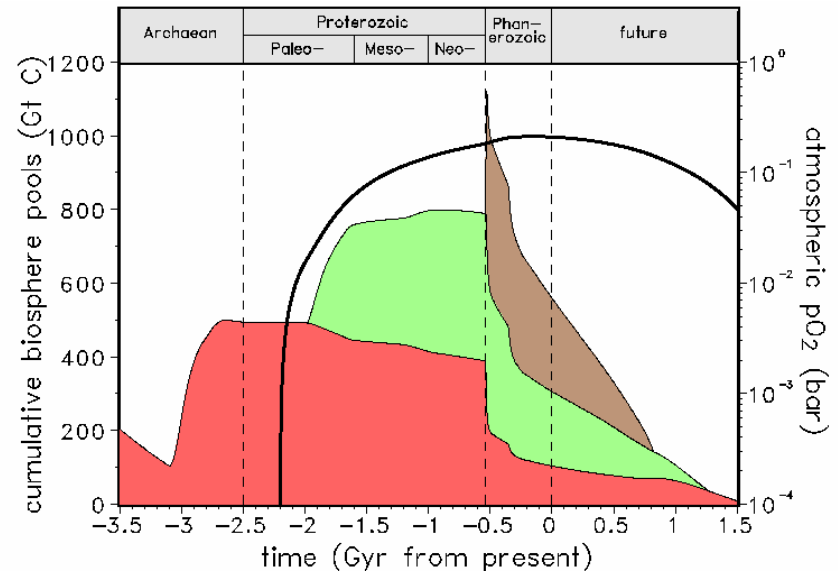
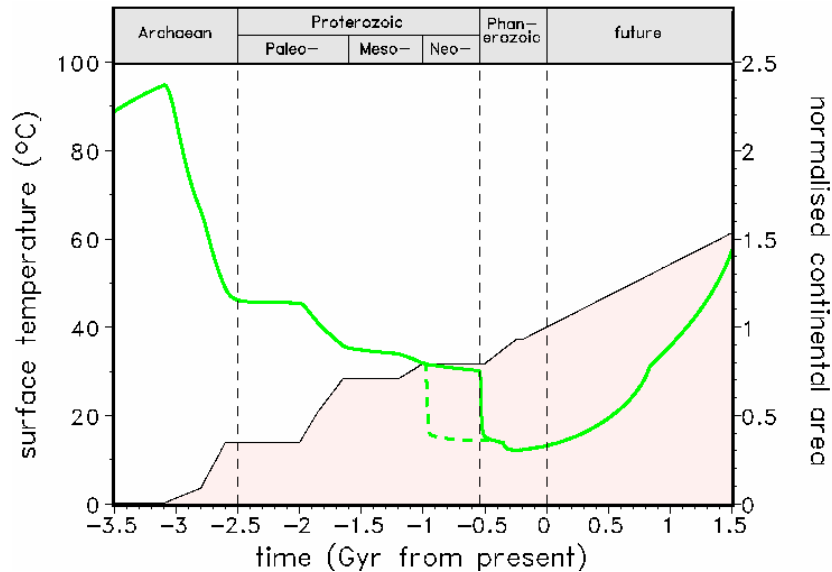
(Lenton and von Bloh, *GRL* 28 (9),1715, 2001)

$$\beta = 1 - \sum_{i=1}^3 \left(1 - \frac{1}{\beta_i} \right) \left(1 - \frac{\Pi_{\text{bio},i}}{\Pi_{\text{bio},1}^*} \right)$$

$$\beta_1 = \beta_2 = 1 \quad \beta_3 = 3.6$$



THE CAMBRIAN EXPLOSION



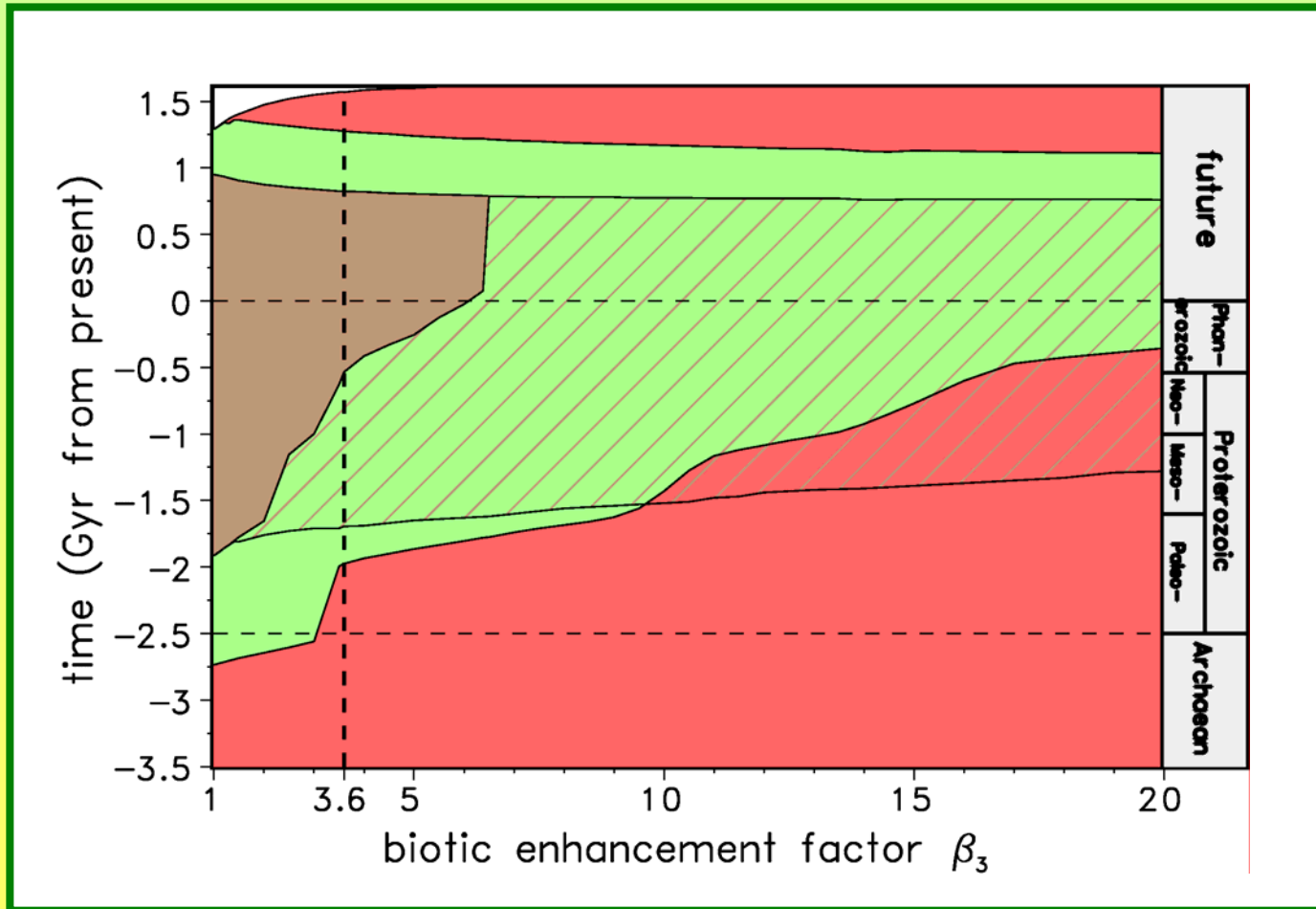
Evolution of global surface temperature.

Evolution of the cumulative biosphere pools for **prokaryotes**, **eucaryotes**, and **complex multicellular life**.

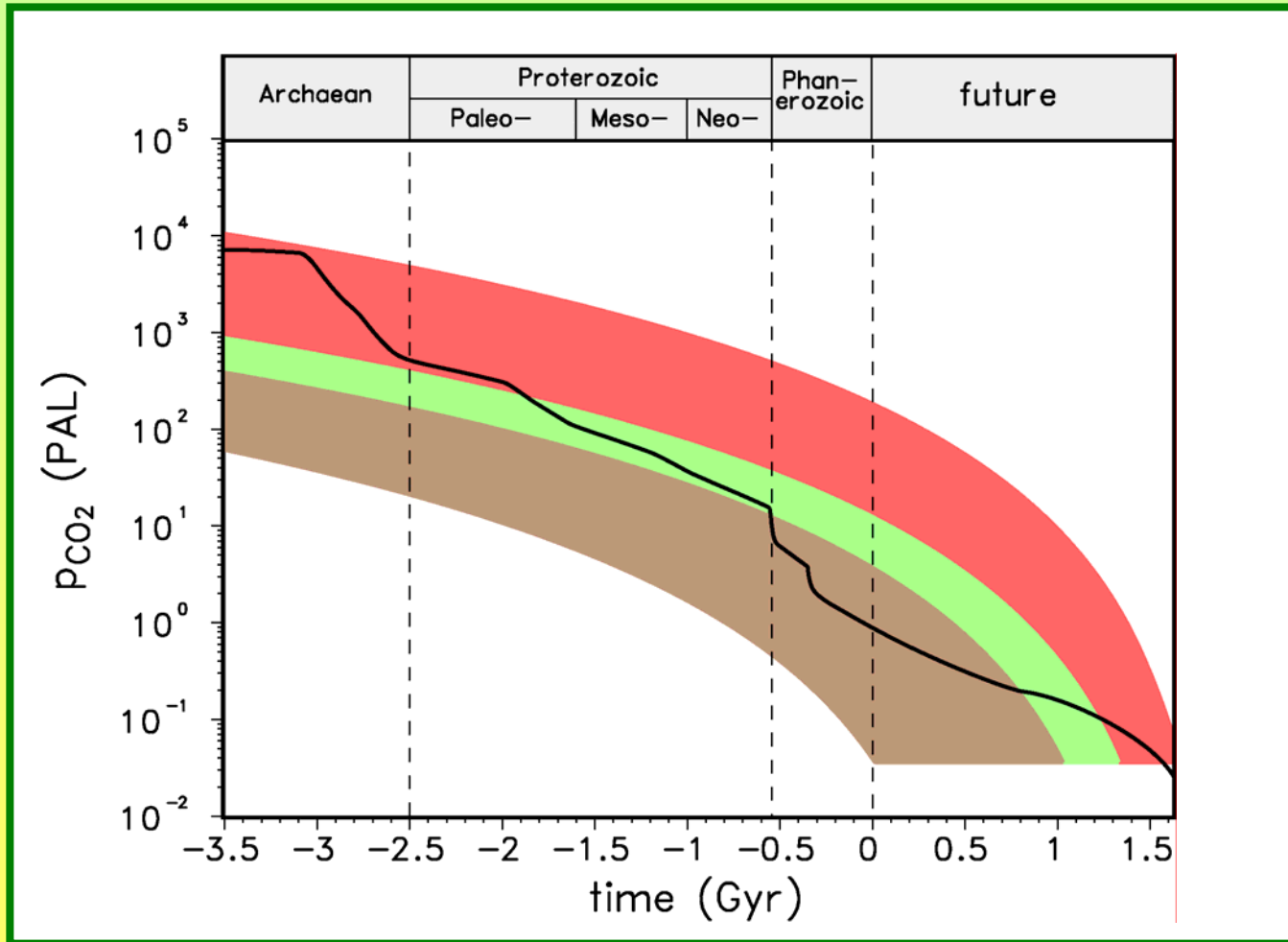
Von Bloh et al. (2003), GRL 30, 1963.



STABILITY DIAGRAM

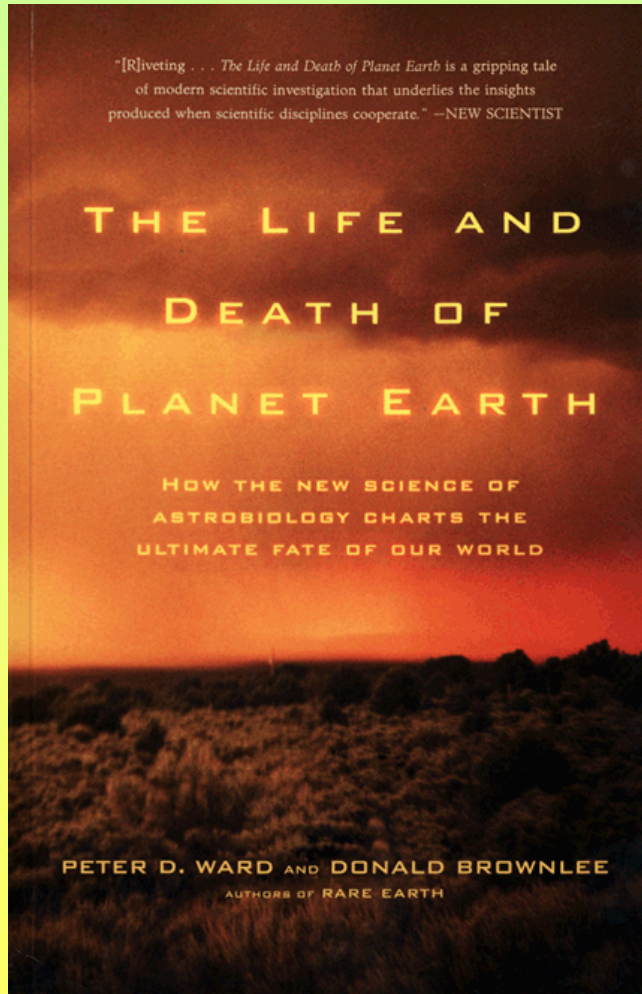


TERRESTRIAL LIFE CORRIDOR (TLC)



Franck et al. (2006), Biogeosciences 3, 85.





- *The life and death of planet Earth*
Peter Ward & Donald Brownlee 2002

A dramatic landscape featuring dark, jagged mountains in the foreground and a large, glowing red sun in the sky. The sun is the central focus, casting a strong red light over the scene. The mountains are silhouetted against the bright red background, creating a stark contrast. The overall atmosphere is one of mystery and potential danger.

THE END ?

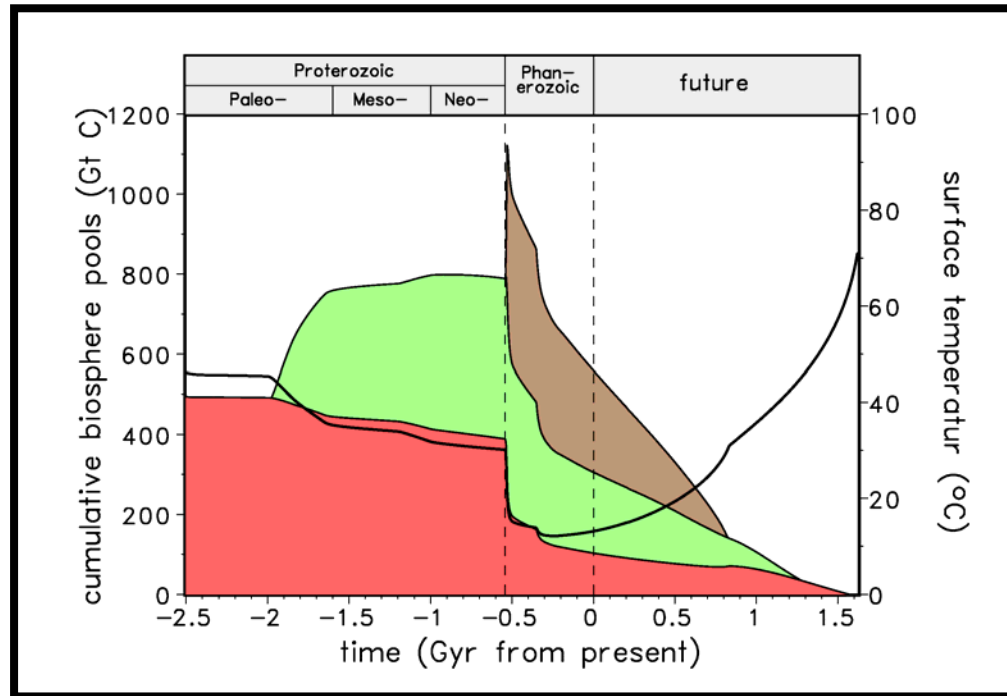
Planetary Engineering

Definition: Planetary engineering is the application of *technology* for the purpose of influencing the global properties of a planet (Fogg, 1995). It is a generalization of the *geoengineering* approach, which investigates modifications of Earth's environment on a global scale to avoid dangerous developments for humankind.



Here: Extending the life span of complex life forms on Earth

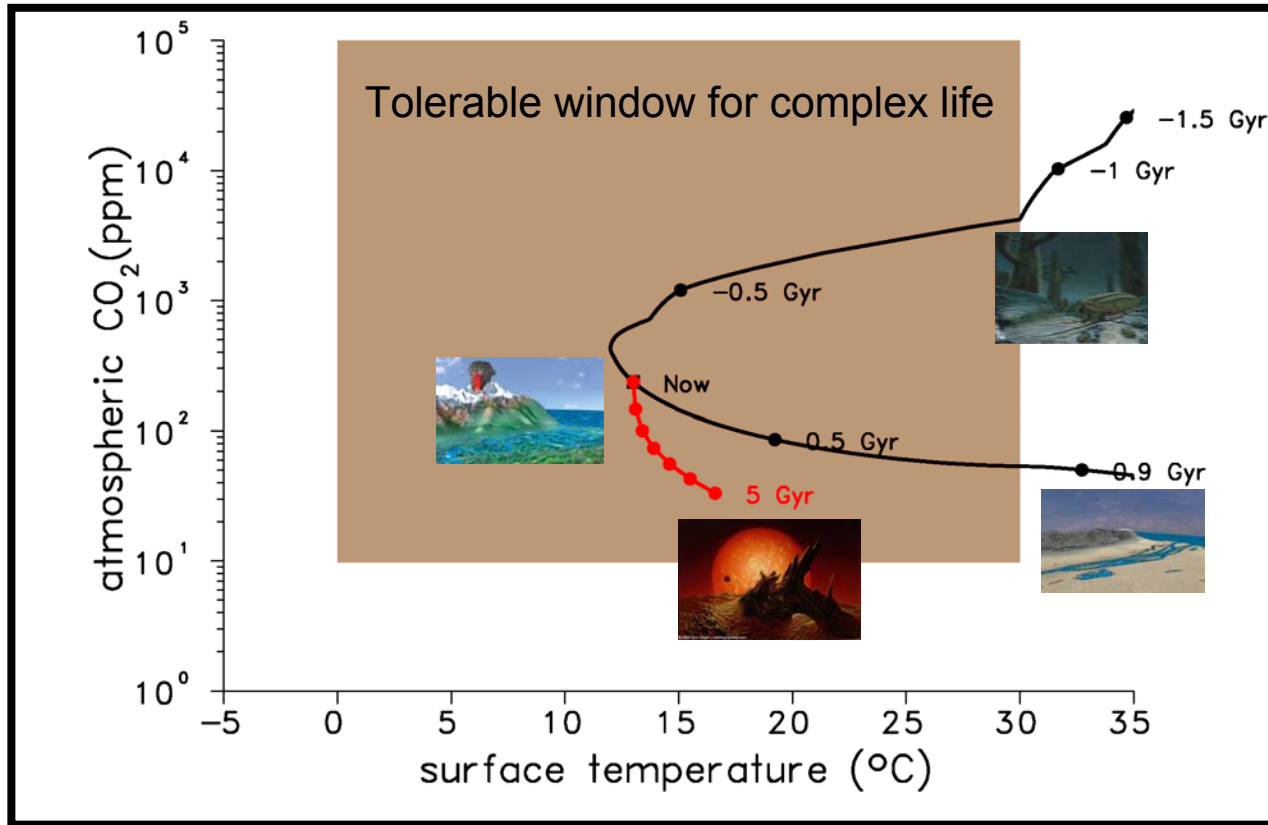
Causes and timing of future biosphere extinctions



S. Franck, C. Bounama, W. von Bloh, *Biogeosciences*, 2006

Biosphere type	Temperature tolerance	CO ₂ tolerance	Biotic enhancement of weathering
Prokaryotes	[5°C, 100°C]	[10, ∞] ppm	1
Eucaryotes (protista)	[2°C, 45°C]	[10, ∞] ppm	1
Complex life	[0°C, 30°C]	[10, ∞] ppm	3.6

Terrestrial Life Corridor (TLC)



TLC defines guardrails for temperature and CO₂



**Special thanks to my co-workers
Werner von Bloh & Christine Bounama**